

The El Niño Southern Oscillation (ENSO) arises from air-sea interactions in the tropical Pacific and affects climate worldwide through atmospheric teleconnections. El Nino is characterized by anomalously warm Sea Surface Temperature (SST) anomalies in the central and by enhanced deep atmospheric convection and westerly wind anomalies in the central equatorial Pacific. While La Niña can broadly be viewed as a mirror image of El Niño, asymmetries between these two ENSO phases have recently become a prominent research interest.

This interest relates to the fact that El Niño events can occasionally reach much larger amplitudes than La Niña events, like in 1982, 1997 and 2015. The most obvious ENSO asymmetry is related to its amplitude, with stronger El Niño than La Niña SST and rainfall anomalies in the eastern Pacific and the nonlinearities (both atmosphere and ocean) are required to explain ENSO asymmetries (An et al. 2020). Here, in the study we focus on the relative roles of atmospheric and asymmetrical SST anomalies forcing to ENSO asymmetrical rainfall response during its peak phase.

**Data and Methods** 

□ The monthly mean fields of SST, precipitation, evaporation and the 3-dimensional specific humidity, zonal and meridional components from the latest version of European Centre for Medium-Range Weather Forecasts (ECMWF) reanalysis (ERA5; Hersbach et al. 2020).

□ The vertically integrated water budget

$$\frac{d < W >}{dt} = < E > - < P > + < V_{tot} > + < I >$$
(1)

$$< V_{tot} > \equiv -\langle \vec{\nabla} \left( \frac{1}{g} \int_{p_s}^{p_t} \vec{v} q dp \right) > = -\vec{\nabla} \left( \frac{1}{g} \int_{p_s}^{p_t} (\langle \vec{v} \rangle \langle q \rangle + \langle \vec{v}'' q'' \rangle) dp \right) \approx V_{lf} + V_{hf}$$
<sup>(2)</sup>

Equation (1) states that low-frequency filtered total atmospheric column water vapour (<W>). The source term associated with surface evaporation (<E>), a sink term precipitation (<P>), and a sink or source term associated with vertically-integrated moisture convergence ( $\langle V_{tot} \rangle$ ) and  $\langle I \rangle$  is the residual. We further decompose  $V_{tot}$  into contributions from intra-seasonal ( $V_{hf}$ ) and lowfrequency  $(V_{lf})$  q and  $\vec{v}$  components.  $V_{lf}$  is the first-order linear function of convergence. Given this relationship,  $V_{lf}$  can be approximated as:

$$V_{lf} \approx V = -m(T)(d-d_0)$$

Here, m(T) is the SST-dependent gross moisture stratification and can be obtained as the slope of various curves displayed in Figure 4.



- ✓ An, S.-I., et al. 2020: ENSO irregularity and asymmetry. In A. Santoso, M. McPhaden & W. Cai (Eds.), El Niño Southern Oscillation in a changing climate (pp. 153–172). John Wiley & Sons.
- ✓ Hersbach, H., and Coauthors, 2020: The ERA5 global reanalysis. Quart. J. Roy. Meteor. Soc., 146, 1999–2049, https://doi.org/10.1002/qj.3803.
- Srinivas, G., et al. 2022: Relative Contributions of Sea Surface Temperature and Atmospheric Nonlinearities to ENSO Asymmetrical Rainfall Response. J. Climate, 35, 3725–3745, https://doi.org/10.1175/JCLI-D-21-0257.1.

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(3)